

Coeurs dynamiques

**FORMATION MODÉLISATION NUMÉRIQUE DE L'OCÉAN ET DE L'ATMOSPHÈRE
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Plan

- **ÉTUDE DE SCHÉMAS TEMPORELS: CAS DES SYSTÈMES**
- **DIFFUSION**
- **SYSTÈMES RAIDES**

Cas des systèmes

Schémas d'advection : cas des systèmes

$$\left\{ \begin{array}{l} \frac{\partial \eta}{\partial t} + u_0 \frac{\partial \eta}{\partial x} + H \frac{\partial u}{\partial x} = 0 \\ \frac{\partial u}{\partial t} + u_0 \frac{\partial u}{\partial x} + g \frac{\partial \eta}{\partial x} = 0 \end{array} \right. \begin{array}{l} \text{SHALLOW WATER} \\ \text{INTERNAL GRAVITY WAVES} \\ \text{LIEN AVEC DISCRÉTISATION PRESSION (TRACEURS)-VITESSES} \end{array}$$

$$u(x, t) = \sum_k U_k(t) e^{ikx}$$

$$U_k(t) = U_k(0) e^{i(kx - \omega t)}, \eta_k(t) = \eta_k(0) e^{i(kx - \omega t)}$$

$$\eta(x, t) = \sum_k \eta_k(t) e^{ikx}$$

RELATION DE DISPERSION : $\omega_{\pm} = k \left(u_0 \pm \sqrt{gH} \right)$

Schémas d'advection : cas des systèmes

$$\left\{ \begin{array}{l} \frac{\partial \eta}{\partial t} + u_0 \frac{\partial \eta}{\partial x} + H \frac{\partial u}{\partial x} = 0 \\ \frac{\partial u}{\partial t} + u_0 \frac{\partial u}{\partial x} + g \frac{\partial \eta}{\partial x} = 0 \end{array} \right. \quad \frac{\partial y^\pm}{\partial t} + \left(u_0 + \sqrt{gH} \right) \frac{\partial y^\pm}{\partial x} = 0, \quad y^\pm = u \pm \frac{g}{H} \eta$$

ÉTUDE DU SCHÉMA TEMPOREL (EXEMPLE : FORWARD-BACKWARD)

$$\left\{ \begin{array}{l} \frac{\eta^{n+1} - \eta^n}{\Delta t} + u_0 \frac{\partial \eta^n}{\partial x} + H \frac{\partial u^n}{\partial x} = 0 \\ \frac{u^{n+1} - u^n}{\Delta t} + u_0 \frac{\partial u^n}{\partial x} + g \frac{\partial \eta^{n+1}}{\partial x} = 0 \end{array} \right. \longrightarrow \begin{pmatrix} \eta_k^{n+1} \\ U_k^{n+1} \end{pmatrix} = A \begin{pmatrix} \eta_k^n \\ U_k^n \end{pmatrix}$$

$$A = \begin{pmatrix} 1 & 0 \\ ikg\Delta t & 1 \end{pmatrix}^{-1} \begin{pmatrix} 1 - iu_0k\Delta t & -iHk\Delta t \\ 0 & 1 - iku_0\Delta t \end{pmatrix}$$

Schémas d'advection : cas des systèmes

$$\begin{pmatrix} \eta_k^{n+1} \\ U_k^{n+1} \end{pmatrix} = A \begin{pmatrix} \eta_k^n \\ U_k^n \end{pmatrix}$$

STABLE SI $\|A\|_2 \leq 1$

VON NEUMANN : ÉTUDE DE LA CONDITION $\rho(A) \leq 1$

(CONDITION NÉCESSAIRE, ÉQUIVALENT SI LA MATRICE EST NORMALE $A^*A = AA^*$)

FORWARD BACKWARD (AVEC $u_0 = 0$)

VALEURS PROPRES DE A : $\lambda_{\pm} = 1 - \frac{1}{2}(\omega\Delta t)^2 \pm i(\omega\Delta t)\sqrt{1 - \frac{1}{4}(\omega\Delta t)^2}$

SI $(\omega\Delta t) \leq 2$, $|\lambda_{\pm}| = 1$ SCHÉMA NEUTRE

Schémas semi-lagrangien : cas des systèmes

SCHÉMA EULÉRIEN (FORWARD - BACKWARD) :

$$\left\{ \begin{array}{l} \frac{\eta^{n+1} - \eta^n}{\Delta t} + u_0 \frac{\partial \eta^n}{\partial x} + H \frac{\partial u^n}{\partial x} = 0 \\ \frac{u^{n+1} - u^n}{\Delta t} + u_0 \frac{\partial u^n}{\partial x} + g \frac{\partial \eta^{n+1}}{\partial x} = 0 \end{array} \right. \quad \text{INSTABLE SI } u_0 > 0$$

$$\left\{ \begin{array}{l} \frac{\partial \eta}{\partial t} + u_0 \frac{\partial \eta}{\partial x} + H \frac{\partial u}{\partial x} = 0 \\ \frac{\partial u}{\partial t} + u_0 \frac{\partial u}{\partial x} + g \frac{\partial \eta}{\partial x} = 0 \end{array} \right.$$

SCHÉMA LAGRANGIEN :

$$\left\{ \begin{array}{l} \frac{\eta^{n+1} - \eta_\star}{\Delta t} + H \frac{\partial u^n}{\partial x} = 0 \\ \frac{u^{n+1} - u_\star}{\Delta t} + g \frac{\partial \eta^{n+1}}{\partial x} = 0 \end{array} \right. \quad \eta_\star = \eta^n(x - u_0 \Delta t), u_\star = u^n(x - u_0 \Delta t)$$

$$\text{CFL: } |(\sqrt{gH})\Delta t| \leq 2$$

SCHÉMA SEMI-IMPPLICIT SEMI-LAGRANGIAN (SISL) :

$$\left\{ \begin{array}{l} \frac{\eta^{n+1} - \eta_\star}{\Delta t} + H \frac{\partial u^{n+1}}{\partial x} = 0 \\ \frac{u^{n+1} - u_\star}{\Delta t} + g \frac{\partial \eta^{n+1}}{\partial x} = 0 \end{array} \right. \quad \text{INCONDITIONNELLEMENT STABLE}$$

$$\eta^{n+1} - gH\Delta t^2 \Delta \eta^{n+1} = \eta_\star - H\Delta t \nabla \cdot u_\star$$

HELMHOLTZ

Plan

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Equation parabolique et schéma implicite

$$\frac{\partial T}{\partial t} = \mathcal{K} \frac{\partial^2 T}{\partial z^2}$$

$$\frac{T^{n+1} - T^n}{\Delta t} = \mathcal{K} \left[(1 - \theta) \frac{\partial^2 T^n}{\partial z^2} + \theta \frac{\partial^2 T^{n+1}}{\partial z^2} \right]$$

$$\theta = \frac{1}{2}$$

CRANK NICHOLSON

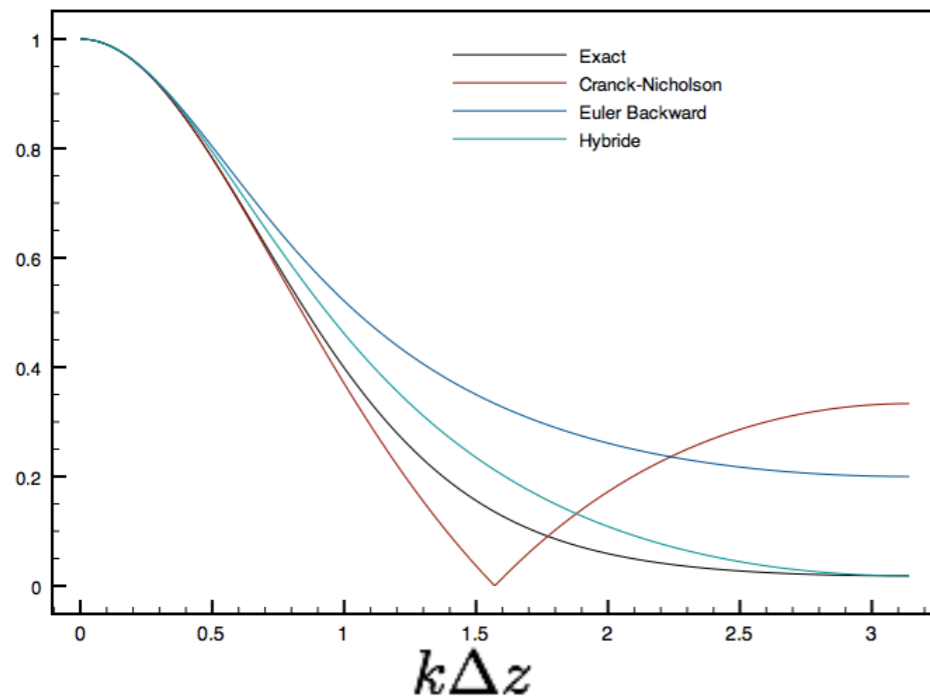
$$\theta = 1$$

EULER BACKWARD

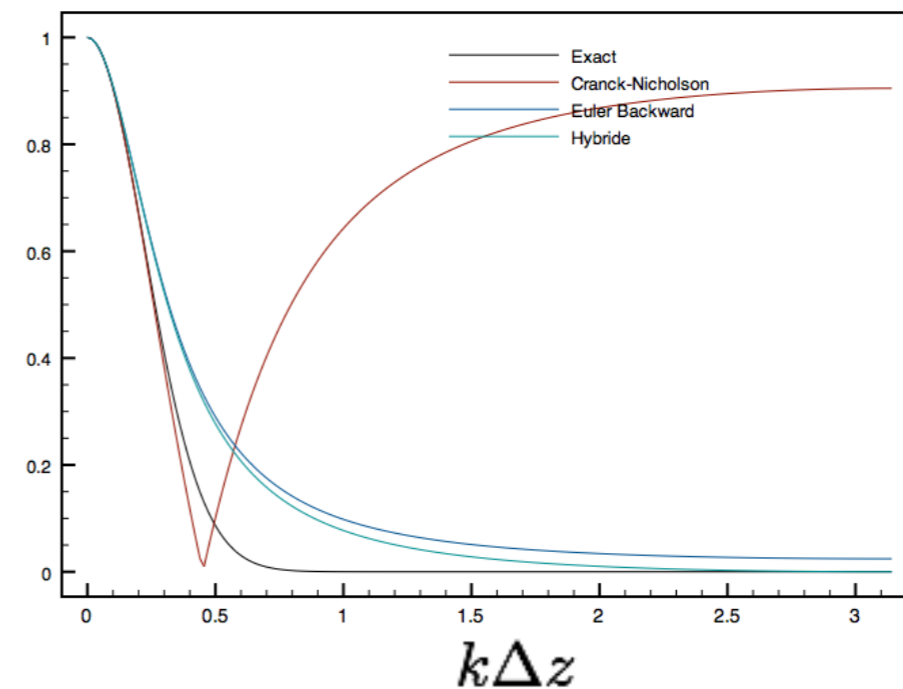
EN EXPLICITE ($\theta = 0$), STABLE SI $\mathcal{K} \frac{\Delta t}{\Delta z^2} \leq \frac{1}{2}$

AMORTISSEMENT

$$\mathcal{K} \frac{\Delta t}{\Delta z^2} = 1$$



$$\mathcal{K} \frac{\Delta t}{\Delta z^2} = 10$$



Equation parabolique et schéma implicite

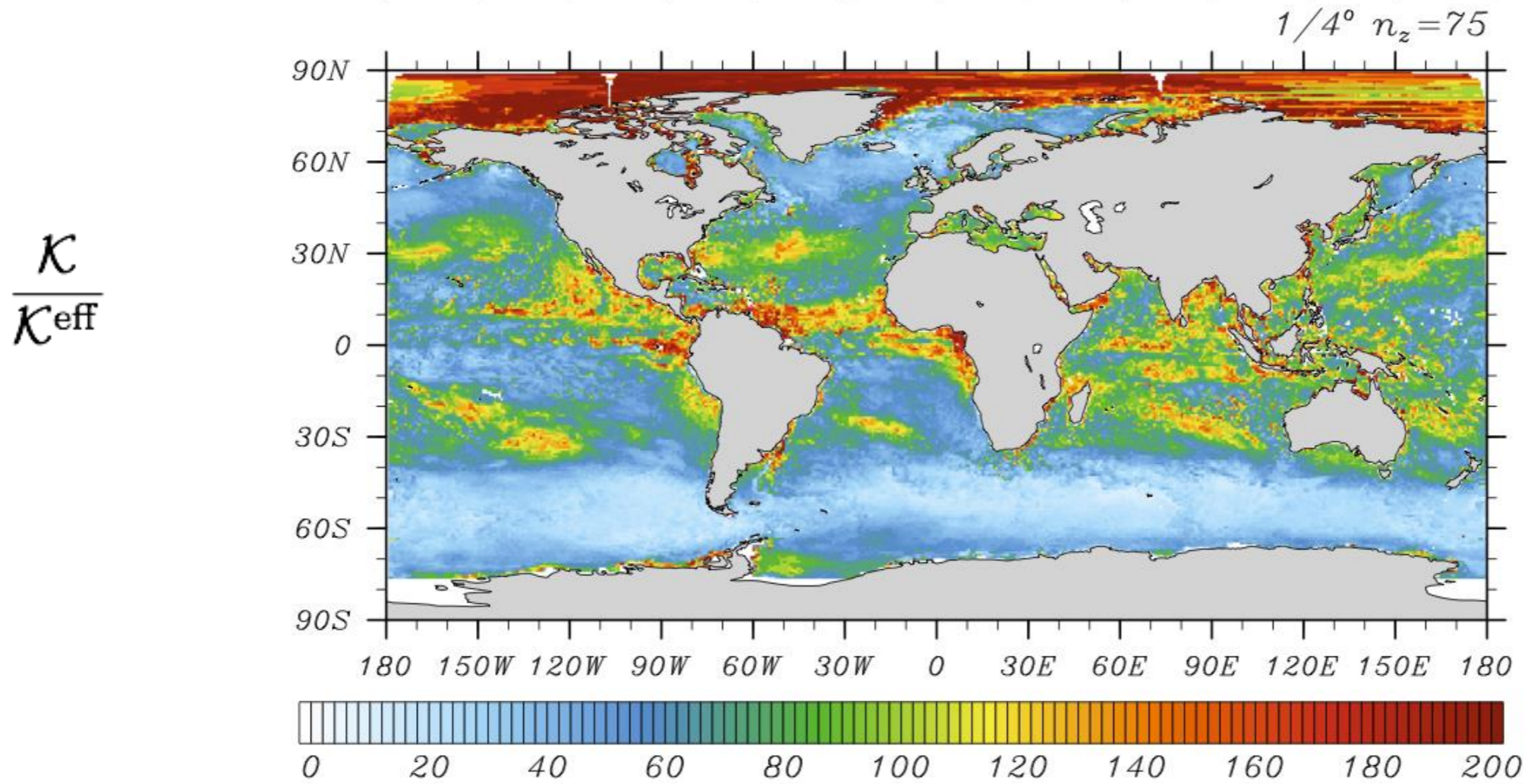


Figure 10: Ratio between the turbulent vertical diffusivity κ and the effective diffusivity κ^{eff} for each water column of the $1/2^\circ$ (top) and the $1/4^\circ$ (bottom) configurations. κ^{eff} is the diffusivity in the continuous equation which would give the same damping as the numerical damping (ideally we should get $\kappa/\kappa^{\text{eff}} = 1$). Areas shaded in yellow and red indicate regions with large numerical errors in the computation of vertical diffusion. The value of $\kappa/\kappa^{\text{eff}}$ is computed using (16) with $\sigma^{(2)} = \bar{\sigma}^{(\text{mld})}$ the averaged value of the vertical parabolic Courant number in the mixed layer and $\theta = 2\pi/N_{\text{mld}}$ with N_{mld} the number of grid points in the mixed layer.

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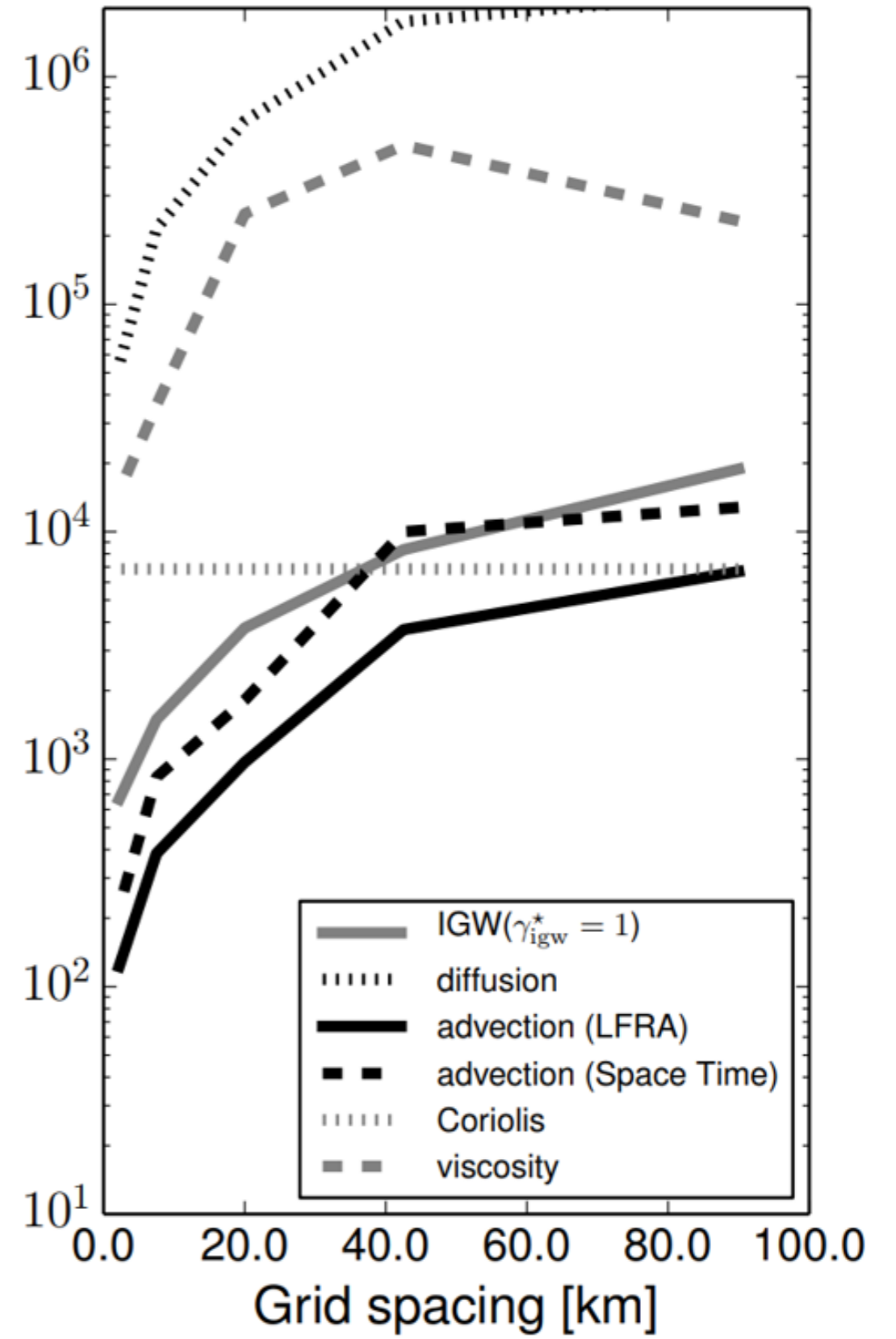
Systemes raides: Methodes de splitting

	Ocean	Atmosphere
Acoustic waves c_s	1500	350
External gravity waves \sqrt{gH}	200	350
Internal gravity waves NH	3	100
Advection U	0.1	30

OCEAN (Boussinesq)

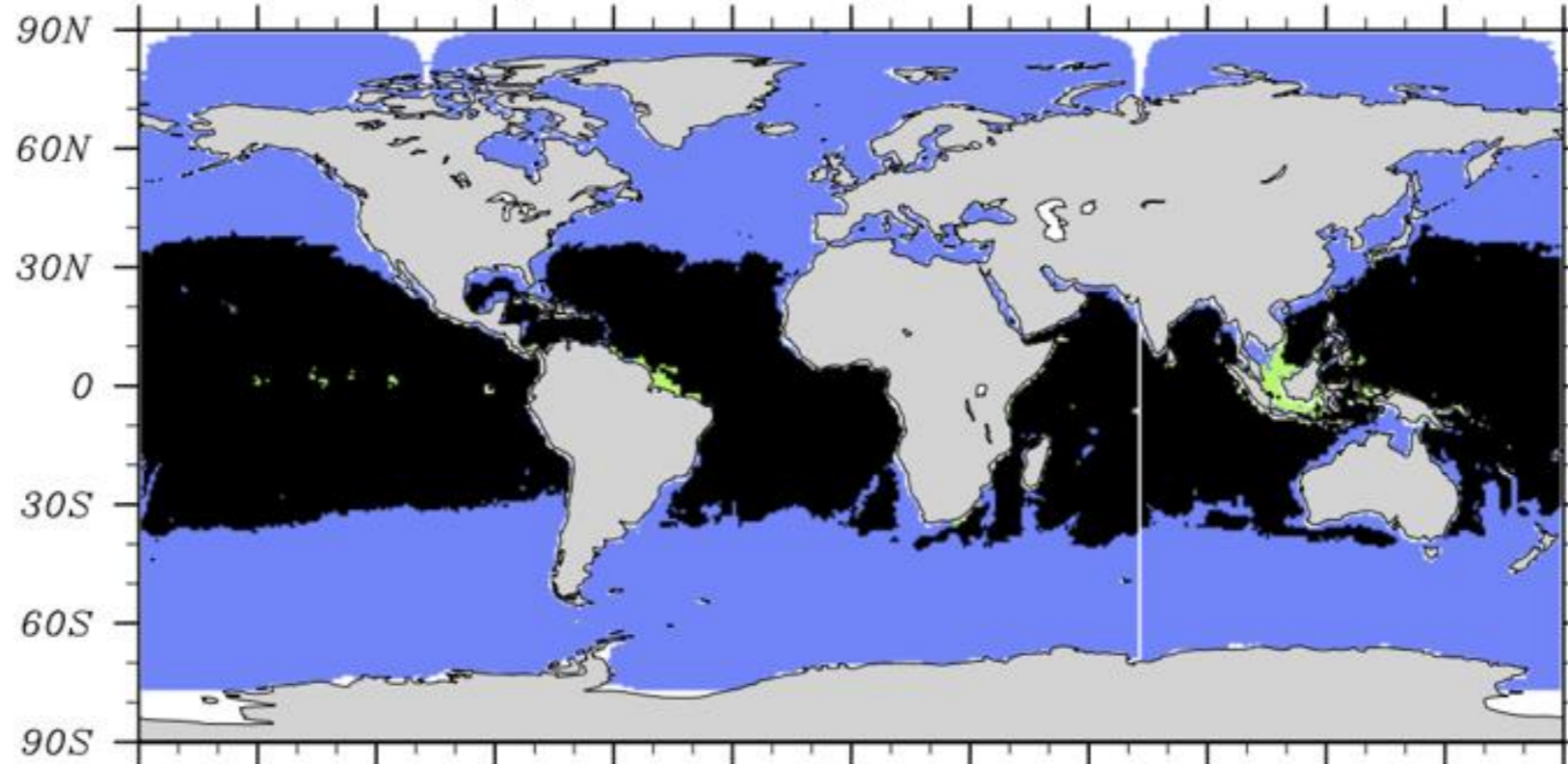
Ocean	
Acoustic waves c_s	1500
External gravity waves \sqrt{gH}	200
Internal gravity waves NH	3
Advection U	0.1

PROCESSUS LIMITANT LE PAS DE TEMPS

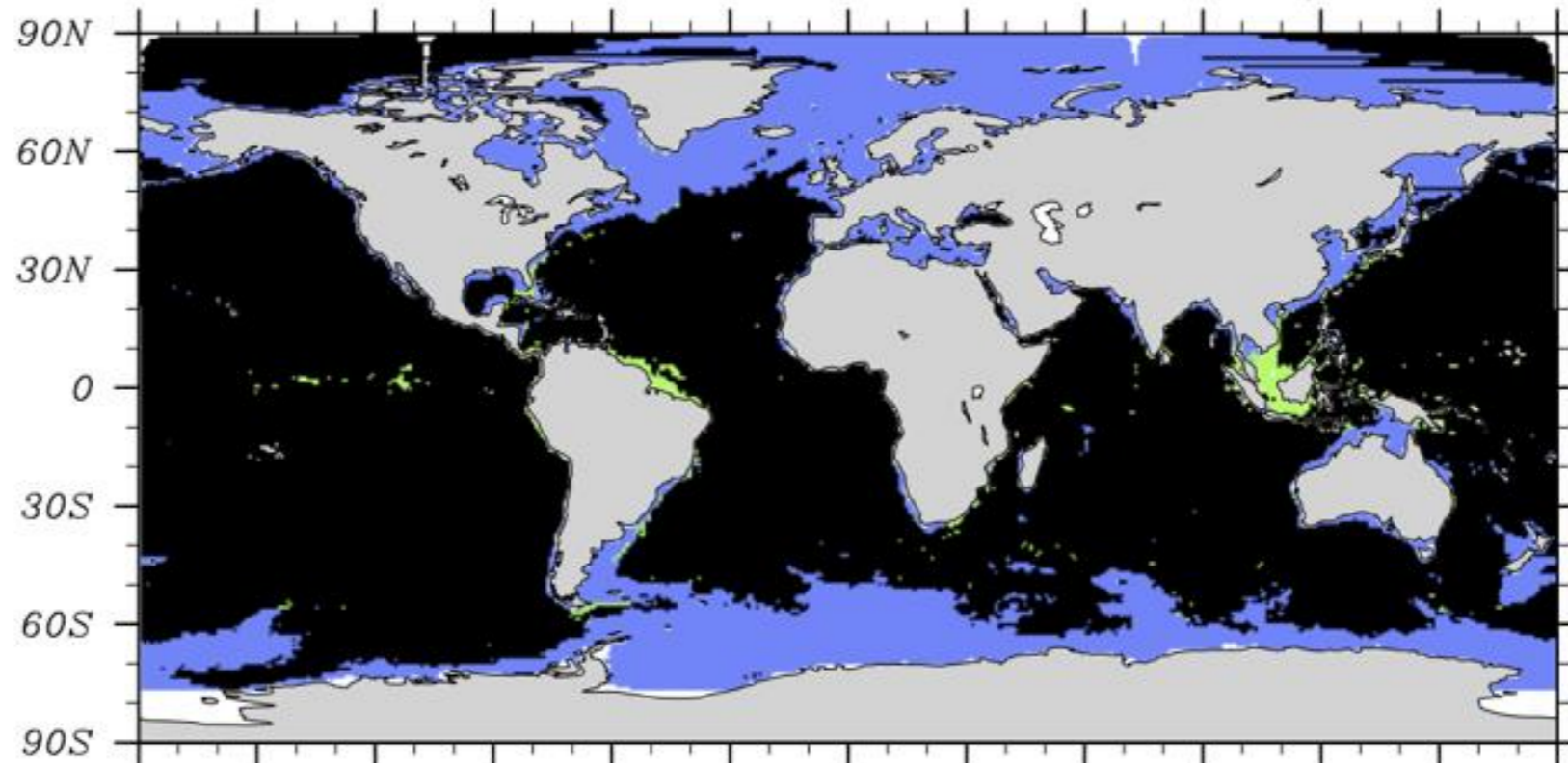


Process limiting the time-step

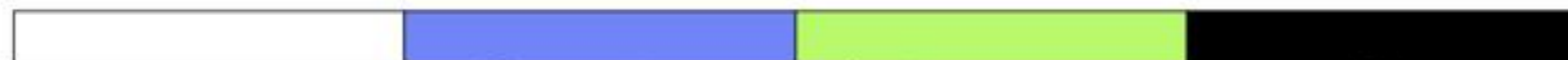
$1/2^\circ n_z=46$



$1/4^\circ n_z=75$



180 150W 120W 90W 60W 30W 0 30E 60E 90E 120E 150E 180

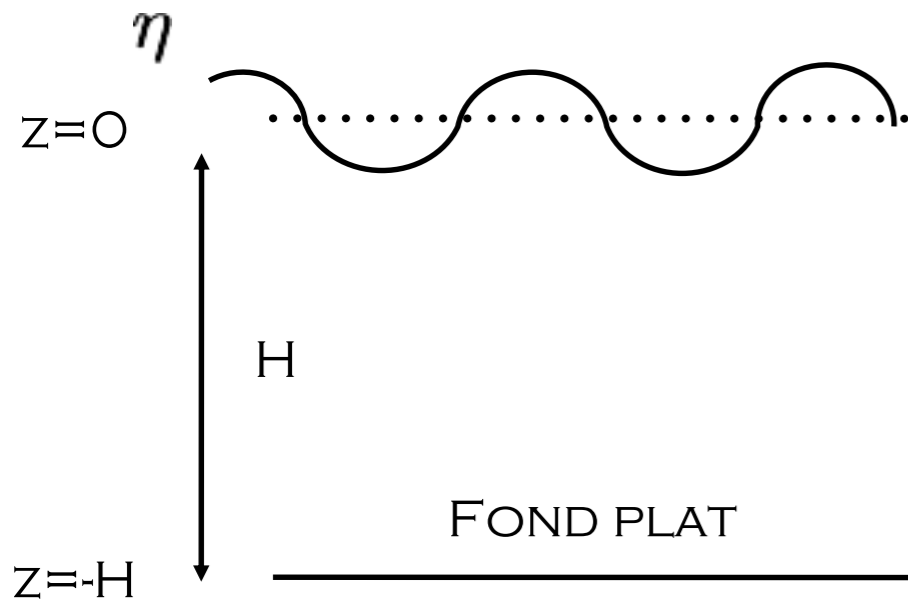


Coriolis

Advection

IGW

1. LINÉARISATION DES ÉQUATIONS PRIMITIVES AUTOUR D'UN ÉTAT MOYEN $(u_0, w_0 = 0, \bar{\rho}(z))$
(2D-VERTICAL)



$$\left\{ \begin{array}{l} \partial_t u' + u_0 \partial_x u' + \frac{1}{\rho_0} \partial_x p' = 0 \\ \partial_x u' + \partial_z w' = 0 \\ \partial_z p' = -\rho' g \\ \partial_t \rho' + u_0 \partial_x \rho' + w' \partial_z \bar{\rho} = 0 \end{array} \right. \quad \left\{ \begin{array}{l} \partial_t \eta = w'(0) \\ w'(-H) = 0 \end{array} \right.$$

2. DÉFINITION D'UNE BASE MODALE ORTHOGONALE MUNIE D'UN PRODUIT SCALAIRE.
SOUS LA FORME D'UN SYSTÈME AUX VALEURS PROPRES DE TYPE STURM-LIOUVILLE

$$\left\{ \begin{array}{l} -\partial_z N^{-2} \partial_z M_q = \lambda_q M_q \\ \frac{dM_q}{dz} \Big|_{z=-H} = 0 \\ \frac{dM_q}{dz} \Big|_{z=0} = -\frac{N^2}{g} M_q(0) \end{array} \right.$$

FRÉQUENCE DE BRUNT-VÄISÄLÄ

$$N^2(z) = -\frac{g}{\rho_0} \partial_z \bar{\rho}(z)$$

$$\langle f, g \rangle = \frac{1}{H} \int_H^0 f g dz$$

$$\langle M_i, M_j \rangle = \delta_{i,j}$$

$$\longrightarrow (M_q, \lambda_q)$$

3. DÉCOMPOSITION DES VARIABLES DANS LA BASE MODALE

$$\begin{aligned} \langle u, M_q \rangle &= u_q & \frac{1}{\rho_0} \langle p, M_q \rangle &= gh_q \\ \sum_q u_q M_q &= u & \sum_q gh_q M_q &= \frac{1}{\rho_0} p & -\rho_0 \sum_q h_q \partial_z M_q &= \rho \end{aligned}$$

4. PROJECTION DES ÉQUATIONS DE CONSERVATION DES MOMENTS, DE DENSITÉ ET DE CONTINUITÉ. OBTENTION D'UNE SOMME ORTHOGONALE DE SYSTÈMES HYPERBOLIQUES (CARACTÉRISTIQUES)

$$c_q = \frac{1}{\sqrt{\lambda_q}} \quad \rightarrow \text{VITESSE DE PHASE}$$

$$\left\{ \begin{array}{l} \partial_t u_q + u_0 \partial_x u_q + g \partial_x h_q = 0 \\ \partial_t h_q + u_0 \partial_x h_q + \frac{c_q^2}{g} \partial_x u_q = 0 \end{array} \right. \iff \left\{ \begin{array}{l} \partial_t \left(u_q + \frac{g}{c_q} h_q \right) + (u_0 + c_q) \partial_x \left(u_q + \frac{g}{c_q} h_q \right) = 0 \\ \partial_t \left(u_q - \frac{g}{c_q} h_q \right) + (u_0 - c_q) \partial_x \left(u_q - \frac{g}{c_q} h_q \right) = 0 \end{array} \right.$$

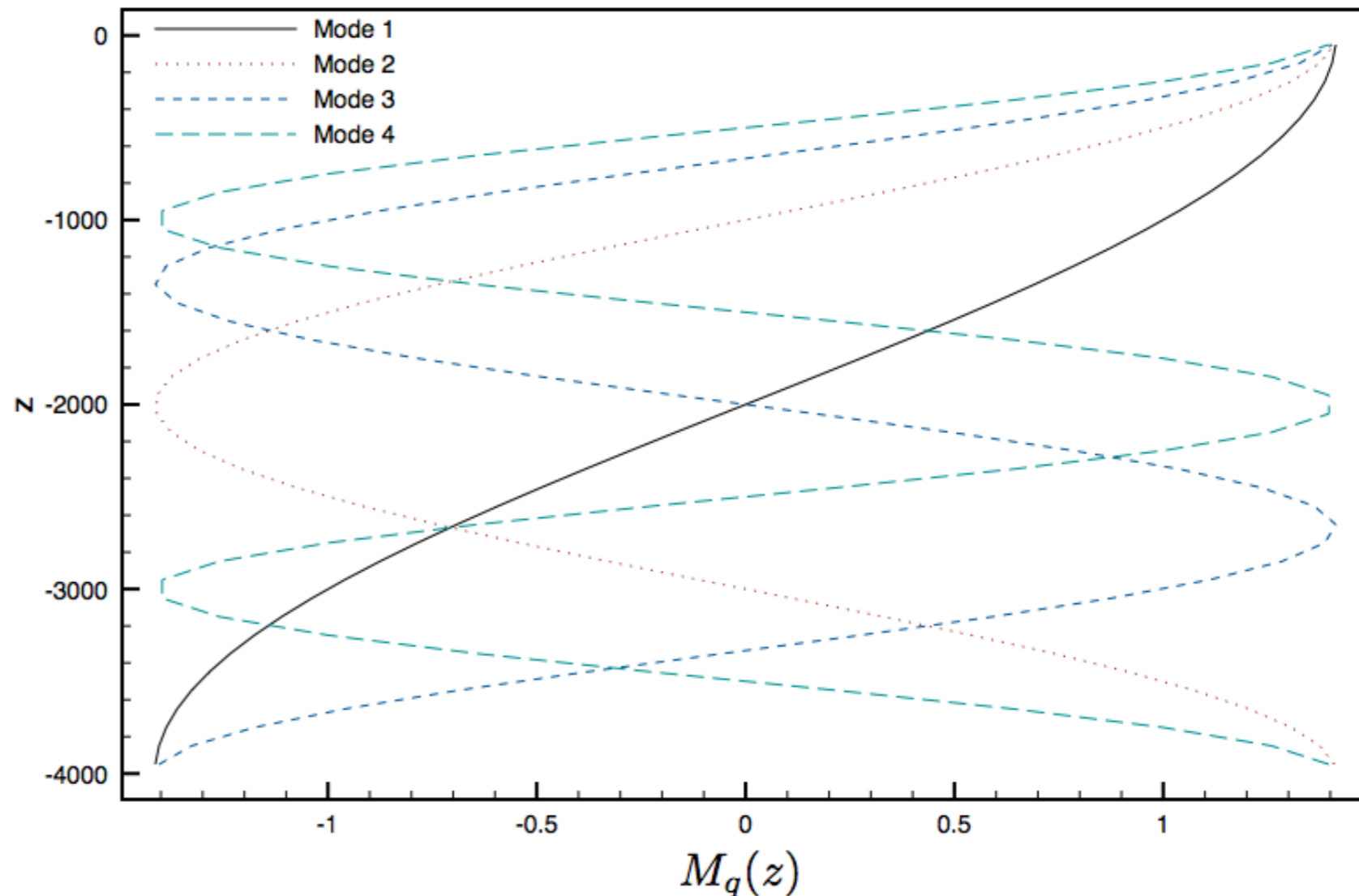
ONDE DE GRAVITÉ :
 SYSTÈME HYPERBOLIQUE COUPLÉ VITESSE-TRACEUR
 INDÉPENDANTE
 VITESSE CONSTANTE : $u_0 \pm c_q$

N CONSTANT : UNE ONDE EXTERNE (BAROTROPE) > DES ONDES INTERNES (BAROCLINES)

$$c_0 \simeq \sqrt{gH} (\simeq 200m/s) \gg c_q \simeq \frac{NH}{q\pi} (c_1 \simeq 3m/s)$$

MODE BAROTROPE QUASI-CONSTANT
 SUR LA VERTICALE.

$$M_0(z) \simeq 1$$



STRUCTURE VERTICALE DES PREMIERS MODES
 BAROCLINES

Barotropic dynamics and time splitting

- THE CFL STABILITY CONDITION ON THE BAROTROPIC MODE LIMITS THE TIME STEP

$$\Delta t_{\text{ext}} < \Delta x / C_{\text{ext}} \text{ where } C_{\text{ext}} = \sqrt{gH} + U_{e_{\text{max}}}$$

$$H = 4000 \text{ m}, C_{\text{ext}} = 200 \text{ m/s (700 km/h)}, \Delta x = 1\text{km}, \Delta t_{\text{ext}} < 5 \text{ s}$$

- BAROCLINIC (INTERNAL) SLOW MODE: (INTERNAL GRAVITY WAVE PHASE SPEED + MAX ADVECTIVE VELOCITY)

$$C_{\text{in}} \approx 2\text{m/s} + U_{i_{\text{max}}} \quad \Delta x = 1\text{km}, \Delta t_{\text{in}} < 8 \text{ mn}$$

- $\Delta t_{\text{in}} / \Delta t_{\text{ext}} \approx 60 - 100!$

SPLITTING BAROTROPE/BAROCLINE

$$c_0 \simeq \sqrt{gH} (\simeq 200 \text{ m/s}^{-1}) \gg c_q \simeq \frac{NH}{q\pi} (c_1 \simeq 3 \text{ m/s}^{-1})$$

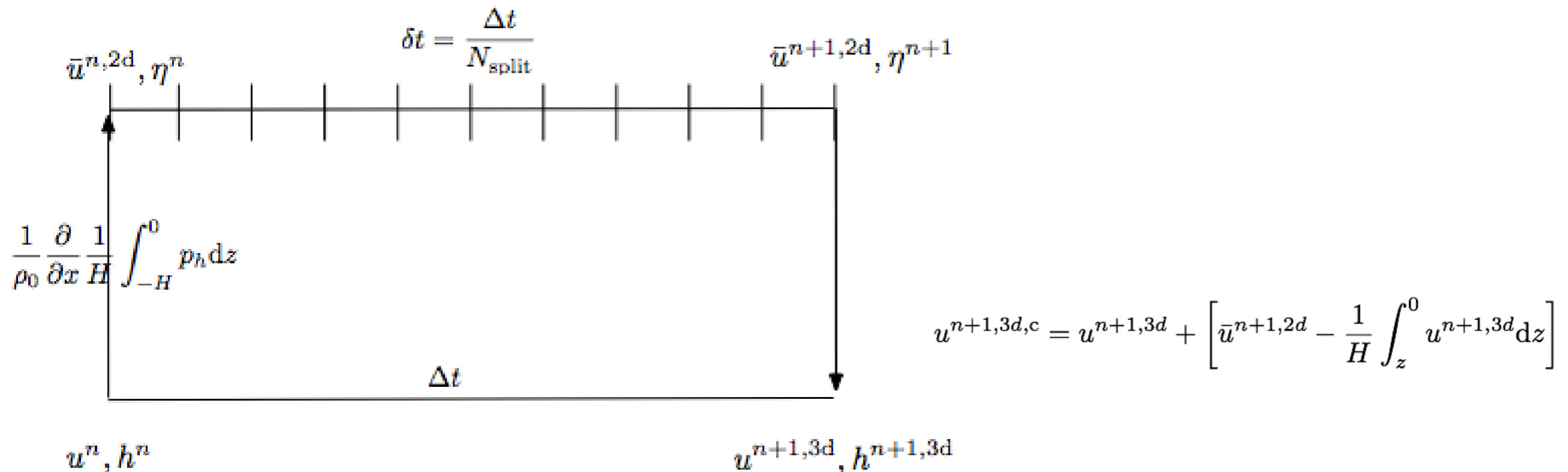
SCHÉMA EXPLICITE → FORTE CONTRAINTE CFL

$$\begin{cases} \frac{\partial \bar{u}}{\partial t} + g \frac{\partial \eta}{\partial x} = -\frac{1}{\rho_0} \frac{\partial}{\partial x} \frac{1}{H} \int_{-H}^0 p_h dz \\ \frac{\partial \eta}{\partial t} + \frac{\partial H \bar{u}}{\partial x} = 0 \end{cases}$$

SYSTÈME 2D : $\delta t \sim \text{CFL BAROTROPE}$

$$\begin{cases} \partial_t \mathbf{u} + (\mathbf{V} \cdot \nabla) \mathbf{u} + \frac{1}{\rho_0} \nabla_h p = 0 \\ \nabla \cdot \mathbf{V} = 0 \\ \partial_z p = -\rho g \\ \partial_t \rho + (\mathbf{V} \cdot \nabla) \rho = 0 \end{cases}$$

SYSTÈME 3D : $\Delta t \sim \text{CFL BAROCLINE}$



Atmosphère

	Ocean	Atmosphere
Acoustic waves c_s	1500	350
External gravity waves \sqrt{gH}	200	350
Internal gravity waves NH	3	100
Advection U	0.1	30

Ondes acoustiques et filtrage temporel

- THE COURANT-FRIEDRICH-LEVY CFL STABILITY CONDITION ON THE ACOUSTIC WAVES LIMITS THE TIME STEP: $c_s = 350\text{m.s}^{-1}$
- EXTERNAL - INTERNAL GRAVITY WAVES : $c_g = 100 - 300\text{m.s}^{-1}$
- ADVECTION : $|U| \leq \frac{c_s}{12}$

$$\left\{ \begin{array}{l} \frac{Du_h}{Dt} + \nabla_h P = 0 \\ \epsilon_{nh} \frac{Dw}{Dt} + \frac{\partial P}{\partial z} = b \\ \epsilon_{nb} \frac{DP}{Dt} + c_s^2 \nabla \cdot u = 0 \\ \frac{Db}{Dt} = 0 \end{array} \right.$$

Filtrage des ondes acoustiques au niveau des équations

$$\left\{ \begin{array}{l} \frac{Du_h}{Dt} + \nabla_h P = 0 \\ \epsilon_{nh} \frac{Dw}{Dt} + \frac{\partial P}{\partial z} = b \\ \epsilon_{nb} \frac{DP}{Dt} + c_s^2 \nabla \cdot u = 0 \\ \frac{Db}{Dt} = 0 \end{array} \right.$$

nb NON BOUSSINESQ

nh NON HYDROSTATIQUE

$\epsilon_{nb}=1, \epsilon_{nh}=1$ PAS DE FILTRAGE

$\epsilon_{nb}=0, \epsilon_{nh}=1$ FILTRAGE ONDES ACOUSTIQUES (REQUIERT UN SOLVEUR 3D POUR LA PRESSION)

Filtrage des ondes acoustiques au niveau des équations

$$\epsilon_{nb}=1, \epsilon_{nh}=1$$

PAS DE FILTRAGE

NON HYDROSTATIQUE, COMPRESSIBLE

$$\left\{ \begin{array}{l} \frac{Du_h}{Dt} + \nabla_h P = 0 \\ \frac{Dw}{Dt} + \frac{\partial P}{\partial z} = b \\ \frac{DP}{Dt} + c_s^2 \nabla \cdot u = 0 \\ \frac{Db}{Dt} = 0 \end{array} \right.$$

Traitement numérique : semi-implicite

$$\left\{ \begin{array}{l} \frac{Du_h}{Dt} + \nabla_h P^{n+1} = 0 \\ \frac{Dw}{Dt} + \frac{\partial P^{n+1}}{\partial z} = b^n \\ \frac{DP}{Dt} + c_s^2 \nabla \cdot u^{n+1} = 0 \\ \frac{Db}{Dt} = 0 \end{array} \right.$$

→ $\Delta P^{n+1} - \frac{1}{(c_s \Delta t)^2} P^{n+1} = G$ HELMHOLTZ

+ ADVECTION SEMI-LAGRANGIEN



SISL

Traitement numérique : time-splitting

$$m = 1, \dots, N \quad \delta t = \frac{\Delta t}{N}$$

TERMES LENTS

$$\left\{ \begin{array}{l} \frac{u^{m+1} - u^m}{\delta t} + \frac{\partial P^m}{\partial x} = -U \frac{\partial u^n}{\partial x} - w^n \frac{\partial U}{\partial z} \\ \frac{w^{m+1} - w^m}{\delta t} + \frac{\partial}{\partial z} \left(\frac{P^m + P^{m+1}}{2} \right) - b^m = -U \frac{\partial w^n}{\partial x} \\ \frac{b^{m+1} - b^m}{\delta t} + N^2 w^{m+1} = -U \frac{\partial b^n}{\partial x} \\ \frac{P^{m+1} - P^m}{\delta t} + c_s^2 \frac{\partial u^{m+1}}{\partial x} + c_s^2 \frac{\partial}{\partial z} \left(\frac{w^{m+1} + w^m}{2} \right) = -U \frac{\partial P^n}{\partial x} \end{array} \right.$$

HEVI : HORIZONTALLY EXPLICIT - VERTICALLY IMPLICIT

Références

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